

BIOSTRATIGRAPHY AND PALEOECOLOGY OF THE ANAH FORMATION (UPPER OLIGOCENE) IN THE SANGAW AREA, SULAIMANIA GOVERNORATE, KURDISTAN REGION, NE IRAQ

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Key words: Anah Formation; Oligocene; Chattian; Biostratigraphy; Paleoecology; Larger benthic foraminifera

ABSTRACT

Biostratigraphy and paleoecology of the Anah Formation are investigated for the first time in the Sangaw area, Sulaimania Governorate, Northeastern Iraq. Forty species of larger foraminifera and skeletal fragments of brachiopod, coral, pelyceps (bivalves), gastropods, bryozoan and algae shell, are described in seven samples collected from one section near the Pungala village. Based on the distribution of the larger benthic foraminifera and other skeletal grains, two assemblage biozones of Upper Oligocene (Chattian) age have been recognized. During the Chattian, the carbonate deposits of the Anah Formation in the studied section are mostly composed of coralline red algae and large flat benthic foraminifera. The biotic associations identified in this study suggest that the carbonate sedimentation of the Anah Formation thrived in tropical to subtropical waters under oligotrophic to mesotrophic middle ramp environment with normal seawater salinity to a higher salinity inner ramp setting and at water depths that ranged from 40 to 80 m.

الطباقية الأحيائية والبيئة القديمة لتكوين عنة (أوليغوسين اعلى) في منطقة سنكاو، محافظة السليمانية، إقليم كردستان، شمال شرق العراق

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المستخلص

يتضمن البحث الحالي دراسة الطباقية الأحيائية والبيئة القديمة لتكوين عنة لأول مرة في منطقة سنكاو، محافظة السليمانية في إقليم كردستان، شمال شرق العراق. أجريت هذه الدراسة على سبعة نماذج من الحجر الكلسي الغني بالمتحجرات جمعت من مقطع واحد قرب قرية بنكالا. تم تشخيص أربعين نوعا من الفورامينيفرا القاعية الكبيرة والمتحجرات الأخرى، وبالاعتماد على هذه الأنواع من الفورامينيفرا والمتحجرات الأخرى، تم التعرف على نطاقين أحيائيين وعلى أساسهما جرى تحديد عمر التكوين في هذا المقطع بالأوليغوسين الأعلى (جائتين). تم استنباط البيئة القديمة من خلال دراسة التنوع الأحيائي لتكوين عنة حيث أن الدلائل تشير إلى أن التكوين ترسب في مياه بحرية استوائية إلى شبه استوائية ضمن بيئة منصبة وسطية ذات ملوحة اعتيادية إلى منصبة داخلية ذات ملوحة أعلى وفي عمق تراوح ما بين 40 – 80 م.

INTRODUCTION

The studied section is located near Pungala village (Mamlaha Anticline), at latitude N: 35° 19' 35" and longitude E: 45° 02' 57" The area is about 30 Km southwest of Cham Chamal town and about 12 Km northwest of Sangaw town, within the Sulaimania

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Governorate (Fig.1). The studied area is characterized by a mountainous terrain; where the mountain ranges trend NW – SE, separated by narrow or wide valleys.

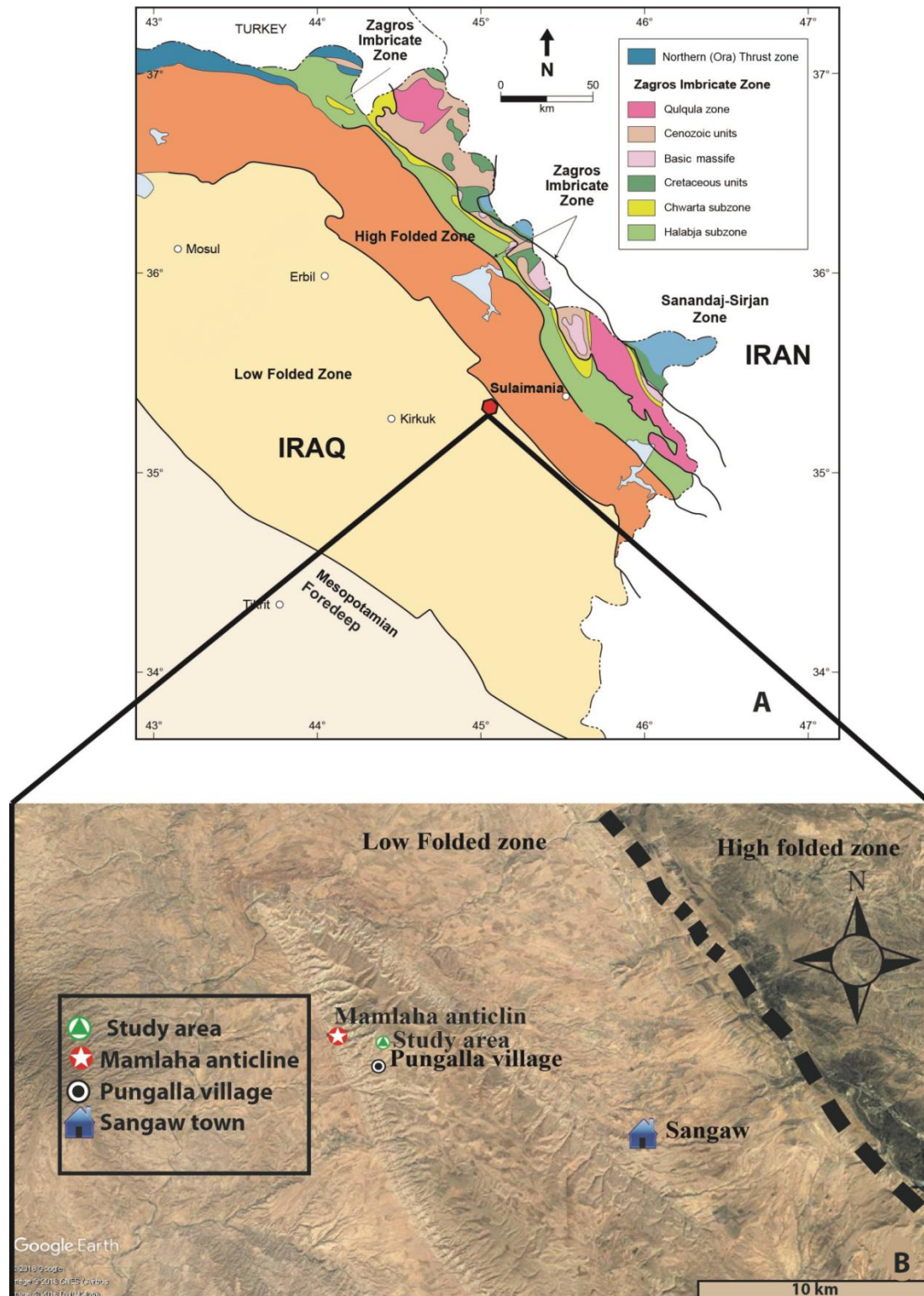


Fig.1: Location map of the studied area. **A)** Tectonic map of Iraq after Al-Kadhimi *et al.* (1996). **B)** Satellite map of the study area taken from Google Earth (2018)

The Anah Formation was introduced by Bellen (1956), the type locality lies 15 Km east of Nahiyah village, west of Anah town, along the Euphrates River and defined by the coordinates: 43° 37' 25" E, 34° 58' 00" N (Bellen *et al.*, 1959). In northern Iraq, the Anah Formation is exposed in Qara Chough and along the Iraqi – Iranian international borders in Shaloradar and Bawgaru anticlines, and as traces along the Bezniyan anticline. According to Al-Banna (1997) in Sissakian and Al-Jubori (2011), the Anah Formation is exposed in the Butmah East anticline. During the regional geological mapping, carried out by Taufiq and Domas (1977) in the Butmah anticline, the base of the Fatha Formation was not found (not exposed); therefore, the formation is not encountered in all the oil wells drilled in the Butmah anticline (I.P.C., 1963). The lower contact is gradational and conformable with the Azkand Formation (Upper Oligocene), whereas, the upper contact with the Euphrates Formation (Miocene) is unconformable (Bellen *et al.*, 1959). The main objective of this paper is to study the biostratigraphy and paleoecology of the Anah Formation in a selected section in NE Iraq, studied for the first time, to determine the age and depositional environment.

▪ Previous Work

- Bellen (1956) described the stratigraphy as the “Main Limestone” of the Kirkuk, Bai Hassan and Qara Chauq Dag structures in Northern Iraq. He reported the following fossils in the type area: algae, anthozoa, bryozoa, echinoidea, mollusca and abundant foraminifera, *Archaias* sp., *Austrotrillina howchini* (Schlumberger), *Borelis pygmaea* Hanzawa, *Heterostegina* cf. *assilinoidea* Blanckenhorn, *Miogypsinoides complanata* (Schlumberger), *Rotalia viennoti* Greig. Numerous undetermined miliolids.
- Ctyroky and Karim (1971) studied the stratigraphy and paleontology of the Oligocene and Miocene strata near Anah town, in the Euphrates valley, and defined the *Miogypsinoides complanata* in the Anah Formation as being of Late Oligocene age.
- Ditmar and the Iraqi – Soviet team (1971) revised Bellen’s classification based on three cycles and divided the Oligocene sedimentary cycle into two sub-cycles. The Lower Cycle comprises the Sheikh Alas, Shurau, Palani and Tarjil formations, and the Upper Cycle consists of the Anah, Azkand, Baba, Bajawan and Ibrahim formations.
- Youkhana and Hradecky (1977) recorded the existence of some Oligocene formations (Shurau, Baba, Bajawan, Azkand and Anah formations) around the Bamu anticline. The thickness of the Oligocene sequence in the recorded area is about 82 m.
- Behnam (1979) studied the stratigraphy and paleontology of the Oligocene – Miocene strata in the Khanaqin area (East Iraq) and described, for the first time from exposed rocks in the area, six Oligocene formations (Shurau formation from the Lower Oligocene, Tarjil, Baba and Bajawan formations from the Middle Oligocene, and Anah and Azkand formations from the Upper Oligocene).
- Buday (1980) confirmed the presence of nine formations in the Kirkuk Group succession.
- Muhammad (1983) studied the biostratigraphy of the formations in the Kirkuk Group in the Kirkuk and Bai Hassan areas.
- Al-Hashimi and Amer (1985) studied the petrographic texture, paleontological character, fossil zonation, facies types and depositional environments of the Oligocene formations as part of their work on the Tertiary microfacies of Iraqi formations.
- Said in Fouad *et al.* (1986) claimed the age of Upper Oligocene for the Anah Formation in the succession cropping out in the Anah area.
- Abid and Sayyab (1989) restudied of the sortidae foraminifera of the Anah Formation and reported an Upper Oligocene – Lower Miocene age.
- Lawa *et al.* (2000) indicated that an Oligocene sequence may be present near the boundary between the High and Low Folded Zones of Iraq.

- Al-Twajjri (2000) suggested that the Anah Formation to be developed during Late Chattian time in a successive superposition.
- Al-Sayyab and Al-Hamdani (1990) studied the biostratigraphy of the Euphrates and Anah formations at Wadi Banat Al-Hassan (W Iraq).
- Abid (1997) and Al-Ghreri (2007) considered the Anah, Azkand and Ibrahim formations to be within the Upper Oligocene – Lower Miocene age.
- Al-Banna (2004), focused on the microfacies analysis of the Oligocene formations in Butmah and Rafan areas Northwest Iraq.
- Sissakian and Mohammed (2007) reviewed stratigraphy of the Iraqi Western Desert that involves the Anah Formation.
- Kharajiany (2008) described the sedimentary facies of the Oligocene rock units at Ashdagh Mountain in Sangaw area, north of Iraq and reported the Anah Formation as of Upper Oligocene age. Al-Banna (2008), studied the Oligocene – Miocene boundary in northern Iraq and revised the previously assigned upper sub-cycle of the Upper Oligocene age (Chattian), suggested by Bellen *et al.* (1959), into an Lower Miocene age (Early Aquitanian).
- Khanaqa *et al.* (2009) studied the lithostratigraphy of a Late Oligocene – Miocene succession in south of Sulaimania city, NE Iraq.
- Al-Banna *et al.* (2010) studied the Oligocene stratigraphy in the Sinjar Basin, northwestern Iraq.
- Karim (2010) modified the time-expanded stratigraphic column of NE Iraq during the Cretaceous and Tertiary times.
- Ghafur (2012) studied the sedimentology and reservoir characteristics of the Oligocene – Early Miocene carbonates (Kirkuk Group) in southern Kurdistan Region.
- Karim *et al.* (2012) studied facies analysis of the Oligocene succession in Sharwaldir anticline, NE of Kalar town, NE Iraq.
- Karim *et al.* (2014) studied the stratigraphy of the Oligocene – Lower Miocene Formations in the Sinjar area, NW Iraq.
- Kharajiany (2014) studied the occurrence of Lower and Middle Miocene rocks (Euphrates, Dhiban and Jerebe formations) in the Ashdagh Mountain, Sangaw area, Sulaimania vicinity, NE Iraq.
- Kharajiany *et al.* (2014) studied the Oligocene and Miocene rock units in Mamlaha anticline, Chamchamal town, Sulaimania Governorate.
- Ghafor and Karim (2014) studied the biostratigraphy of the Oligocene succession in the High Folded Zone, Sulaimania Governorate.
- Al-Qayim *et al.* (2015), studied microfacis and sequence stratigraphy of the Oligocene – Miocene sequence at Golan Mountain.
- Lawa and Ghafur (2015) studied sequence stratigraphy and biostratigraphy of the prolific Late Eocene, Oligocene and Early Miocene carbonates from the Zagros Fold-Thrust Belt in the Kurdistan Region.

SAMPLES AND METHODS

This study is based on seven samples collected from 7 m thick outcrop section of the Anah Formation in the Mamlaha anticline, near Pungala village (Fig.2). The samples are collected on the basis of lithology change and type of fossil content. The lithology of the sequence is described based on field observations. Most of the samples contain well preserved and abundant benthic foraminifera, coral, algae and other skeletal grains of microfossils. The laboratory work included preparation of two thin sections for each sample. The thin sections were examined using polarized microscope to identify different types of microfossils. A

digital camera was used to photo the microfossils in thin sections. The first part of this paper is assigned to biostratigraphic analysis of the Anah Formation, based on benthic foraminifera described by Laursen *et al.* (2009) study. The second part is focused on paleoecological interpretations, such as salinity, nutrients, and depth, based on the studies of many workers (e.g., Geel, 2000; Romero *et al.*, 2002; Brandano *et al.*, 2009 and Flügel, 2010).

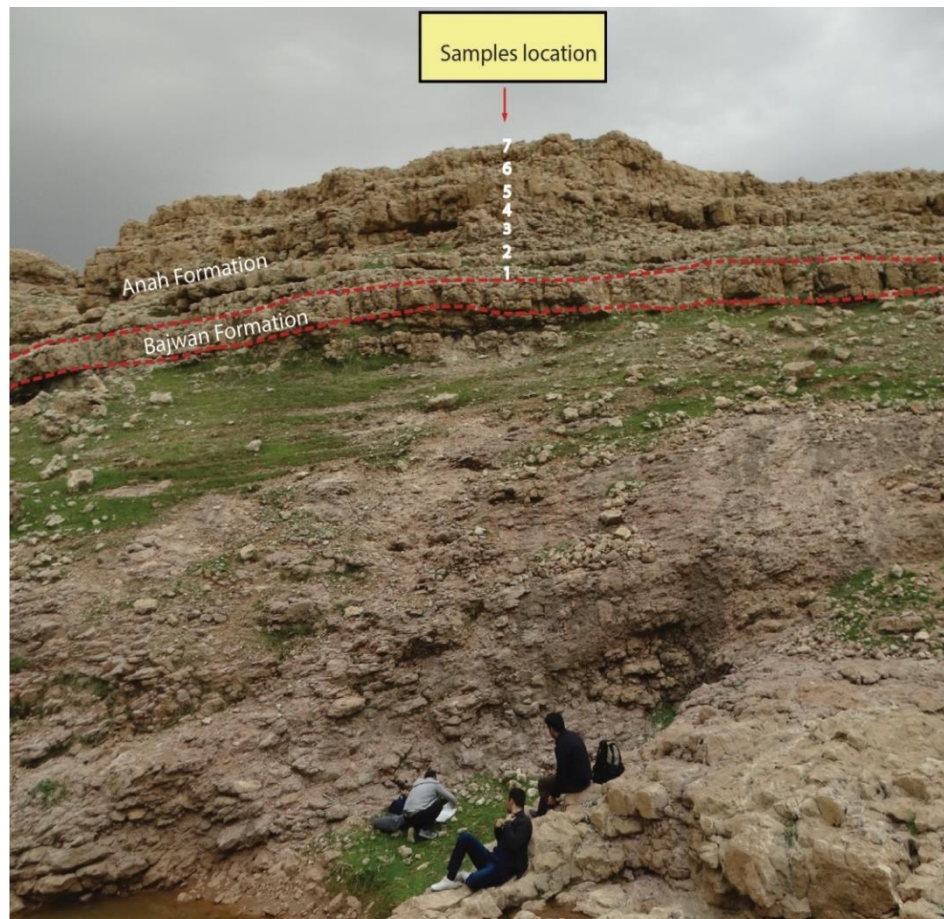


Fig.2: Sampling location showing the Anah Formation and its lower contact with the Bajwan Formation at the Pungala village, Sulaimania Governorate

GEOLOGICAL SETTING AND LITHOSTRATIGRAPHY

Mamlaha anticline locates between southwest of Sulaimani city and the southeast of Kirkuk city, 30 Km SW Chamchamal town, and 13 Km NE Sangaw town. Tectonically the study area is located within the Low Folded Zone. The studied section is situated on the southwestern limb of the Mamlaha anticline which is part of three enechelon anticlines (Ashdagh, Mamlaha and Qara Wais). The Mamlaha anticline is smaller than the other two anticlines, and its limbs have gentle slopes and trends NW – SE like other anticlines in the area. The Ashdagh anticline is located southeast of Mamlaha anticline and Qara Wais anticline is located northeast of the Mamlaha anticline (Fig.3). The studied section is comprised of limestone beds with white thickly bedded highly jointed and fractured coralline limestone. The lower contact is with thin to intermediate red algal limestone of the Bajwan Formation and the upper contact is with the milky thick bedded highly jointed fossiliferous limestone of the Euphrates Formation. The thickness of the Anah Formation in the studied section is about 7 m (Fig.4).

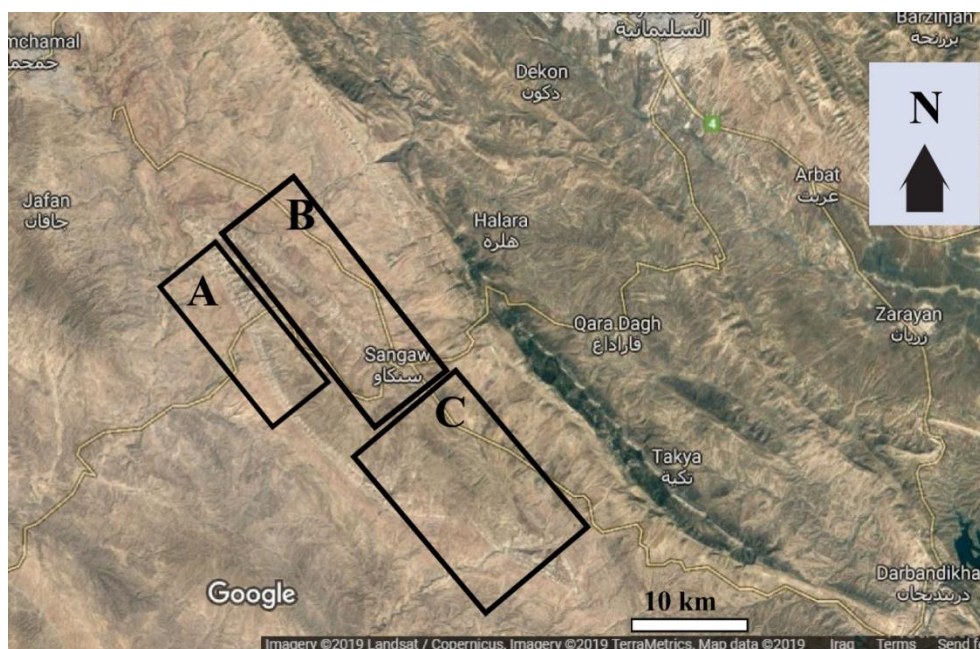


Fig.3: Satellite map of the study area taken from Google Earth (2019) Showing. A) Mamlaha anticline, B) Qara Wais anticline, C) Ashdagh anticline

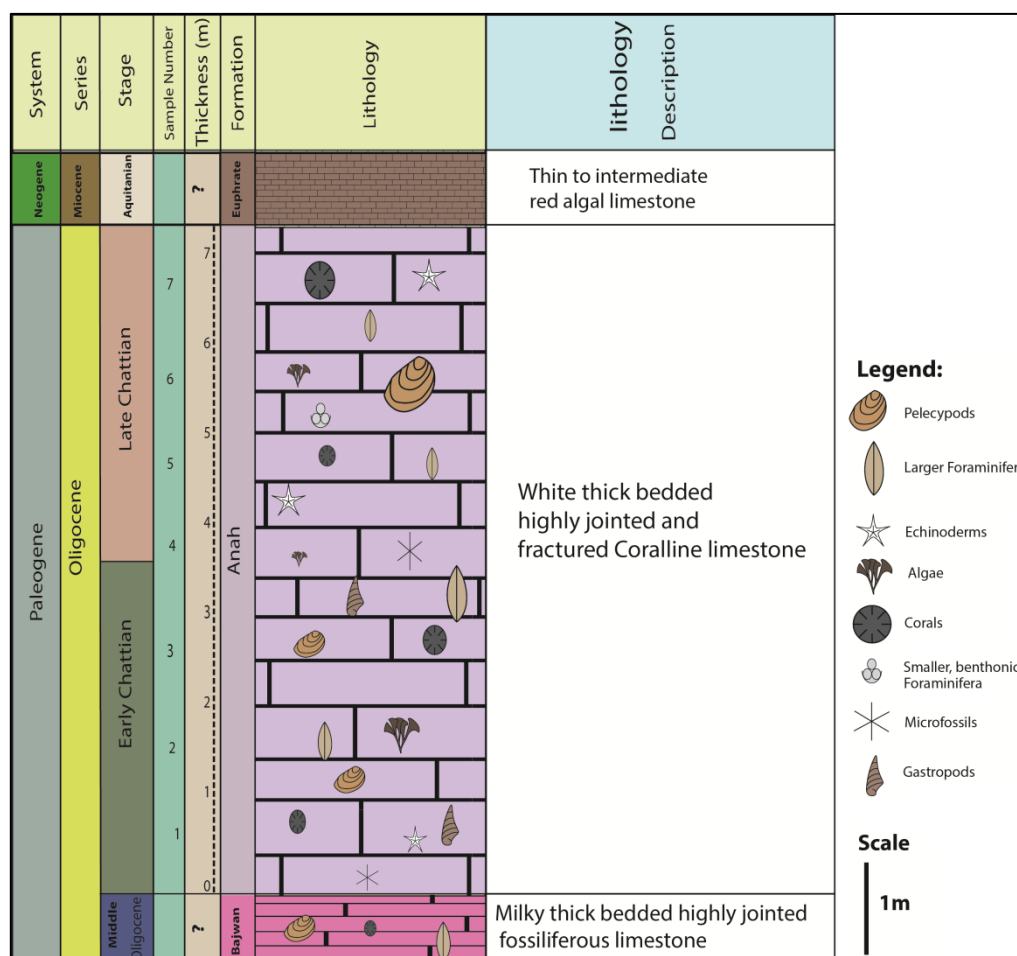


Fig.4: Lithostratigraphic column of the studied section

RESULTS AND DISCUSSION

▪ Biostratigraphy

Forty species from the Late Oligocene foraminiferal assemblages and skeletal fragments of brachiopod layer grain, coral, pelecypods-bivalves, gastropoda shell, bryozoan shell and algae are recognized and analyzed to establish the biostratigraphic framework of the Anah Formation in the studied section (Fig.5). Two assemblage zones are recognized in this study:

– **Assemblage Zone I:** The thickness of this zone is 3 m, recorded from the lower part of the studied section. The most diagnostic species identified include: *Austrollina howchini*; *Quinloculina* sp. *Austrotrillina apaucialveolata*; *Austrotrillina* sp.; *Valvulina*, sp.; *Amphistegina lessonii*; *Peneroplis thomasi*; *-Peneroplis evolutus*; *Textularia* sp.; *Triloculina* sp.; *Triloculina trigonula*; *Triloculina tricarinata*; *Operculina complanata*; *Operculina* sp.; *Nummulites vascus*; *Nummulites* cf. *bouillei*; *Austrotrillina brunni*; *Austrotrillina* sp.; *Spiroclypeus blankenhorni*; *Spiroloculina cylindracea*; *Peneroplis* cf. *farsensis*; *Elphidium* sp.; *Neorotalia viennoti*; *Dendritina rangi*, with other different fossils such as brachiopoda layer grain-large brachiopod grain, coral, pelecypods, gastropoda, echinodermata, bryozoan and algae (Figs.6 – 8). This assemblage is correlated with the *Lepidocyclina – Operculina – Ditrupa* Assemblage Zone of Laursen *et al.* (2009) and Van Buchem *et al.* (2010), and to the SBZ 22 of Cahuzac and Poignant (1997) and to the Assemblage Zone 1 of Taheri *et al.* (2017). Based on Laursen *et al.* (2009) this Assemblage Zone is attributed to the Rupelian-Chattian age. However, Ehrenberg *et al.* (2007), believes that the presence of *Spiroclypeus blankenhorni* indicates Early Chattian age, so the age of this part of the sequence should be limited to Lower Chattian.

– **Assemblage Zone II:** The thickness of this zone is 4m, recorded from the upper part of the studied section. The most diagnostic species include (Figs.6 – 8). *Archaias kirkukensis*; *Risananeiza pustulosa*; *Archaias asmaricus*; *Archaias hensonis*, *Archaias* sp.; *Austrollina howchini*; *Quinloculina* sp.; *Peneroplis thomasi*; *Peneroplis evolutus*; *Triloculina* sp.; *Triloculina trigonula*; *Triloculina tricarinata*; *Archaias operculiniformis*; *Austrotrillina apaucialveolata*; *Austrotrillina* sp.; *Meandropsina anahensis*; *Meandropsina* sp.; *Valvulina* sp.; *Amphistegina lesson*; *Textularia* sp.; *Triloculina* sp.; *Triloculina trigonula*; *Triloculina tricarinata*, *Operculina complanata*; *Operculina* sp.; *Austrotrillina brunni*; *Pyrgo* sp., *Risananeiza crassaparies* with other different fossils such as brachiopoda layer grain-large brachiopod grain, coral, pelecypod shells, gastropoda shells, echinodermata, bryozoan shell and algae (Figs.6 – 8). The microfauna from this biozone correspond to the *Archaias asmaricus – Archaias hensoni – Miogypsinoides complanatus* Assemblage Zone of Laursen *et al.* (2009) and Van Buchem *et al.* (2010), and to the Assemblage Zone 2 of Taheri *et al.* (2017), and indicate Chattian age. The presence of *Risananeiza crassaparies* indicates Late Chattian age, based on Benedetti and Briguglio (2012) and consequently the age of this zone should be limited to Upper Chattian.

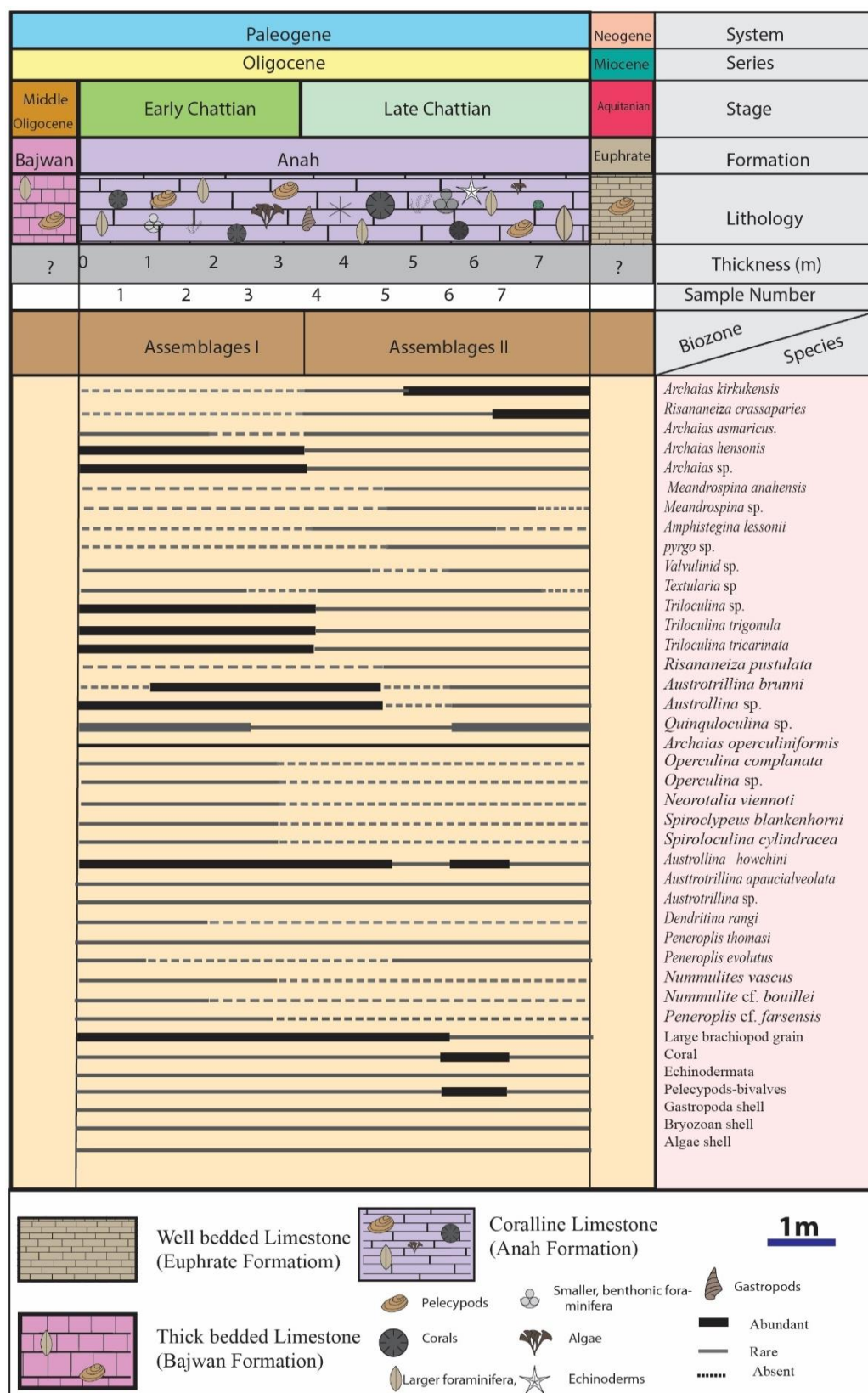


Fig.5: Biostratigraphic column of the Anah Formation in the studied section

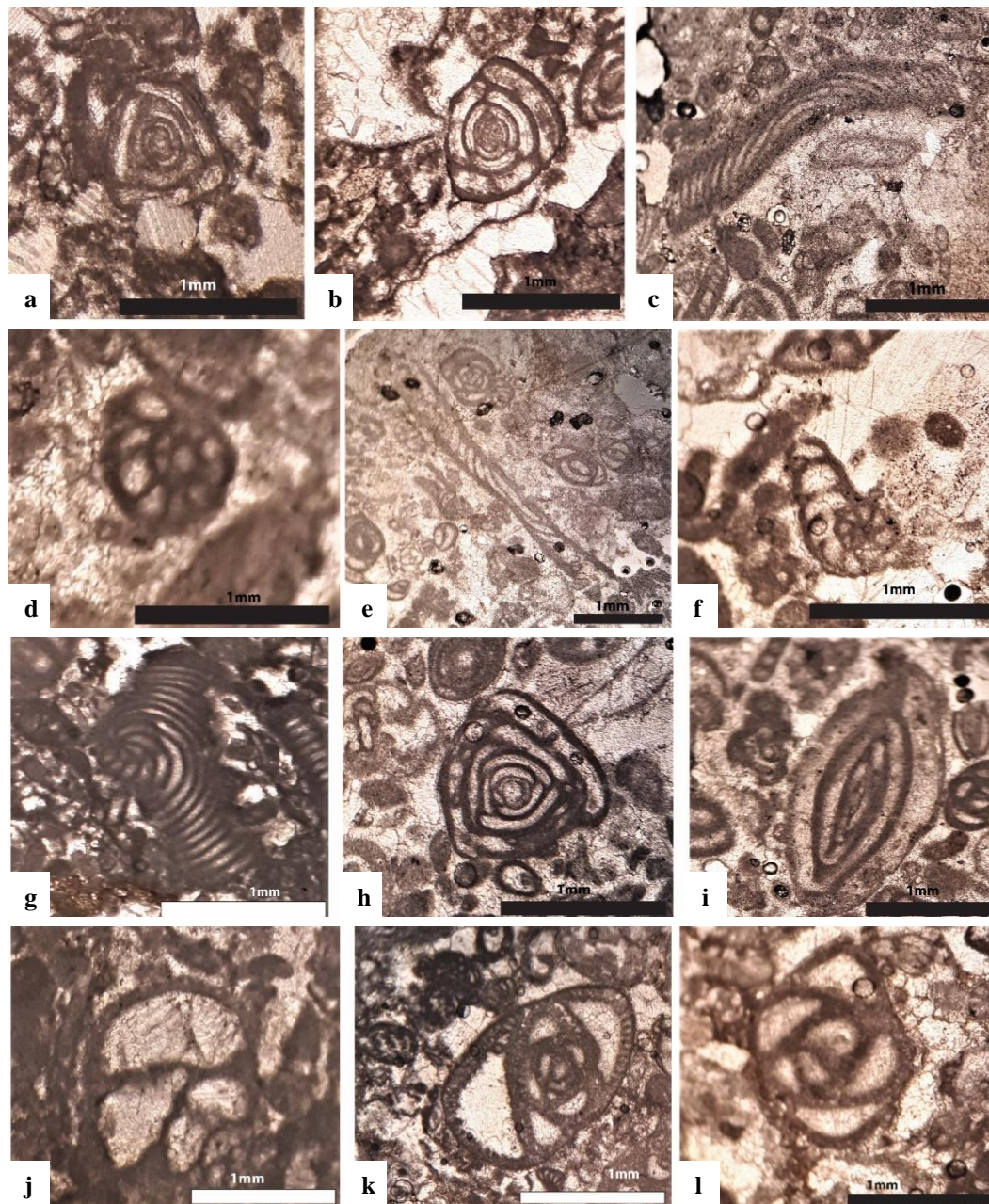


Fig.6: **a)** *Triloculina tricarinata* (d'Orbigny, 1826), sample no. 2; **b)** *Triloculina trigonula*. (Lamarck, 1804), sample no.2; **c)** *Archaias kirkukensis* (Henson, 1950 in Hayward *et al.*, 2018), sample no.3; **d)** *Elphidium* sp. sample no.3; **e)** *Meandropsina anahensis* (Henson, 1950 in Hayward *et al.*, 2018), sample no.3; **f)** *Spiroloculina* cf. *cylindracea* (Henson, 1950 in Hayward *et al.*, 2018), sample no.3; **g)** *Archaias hensoni* (Smout and Eames, 1958), transverse section, sample no.5; **h)** *Triloculina* sp. sample no.2; **i)** *Nummulites* cf. *bouille*, sample no.3; **j)** *Valvulinid* sp., sample no.4; **k)** *Austrotrillina brunni* Marie, transverse section, sample no.4; **l)** *Austrotrillina* sp. sample no.4

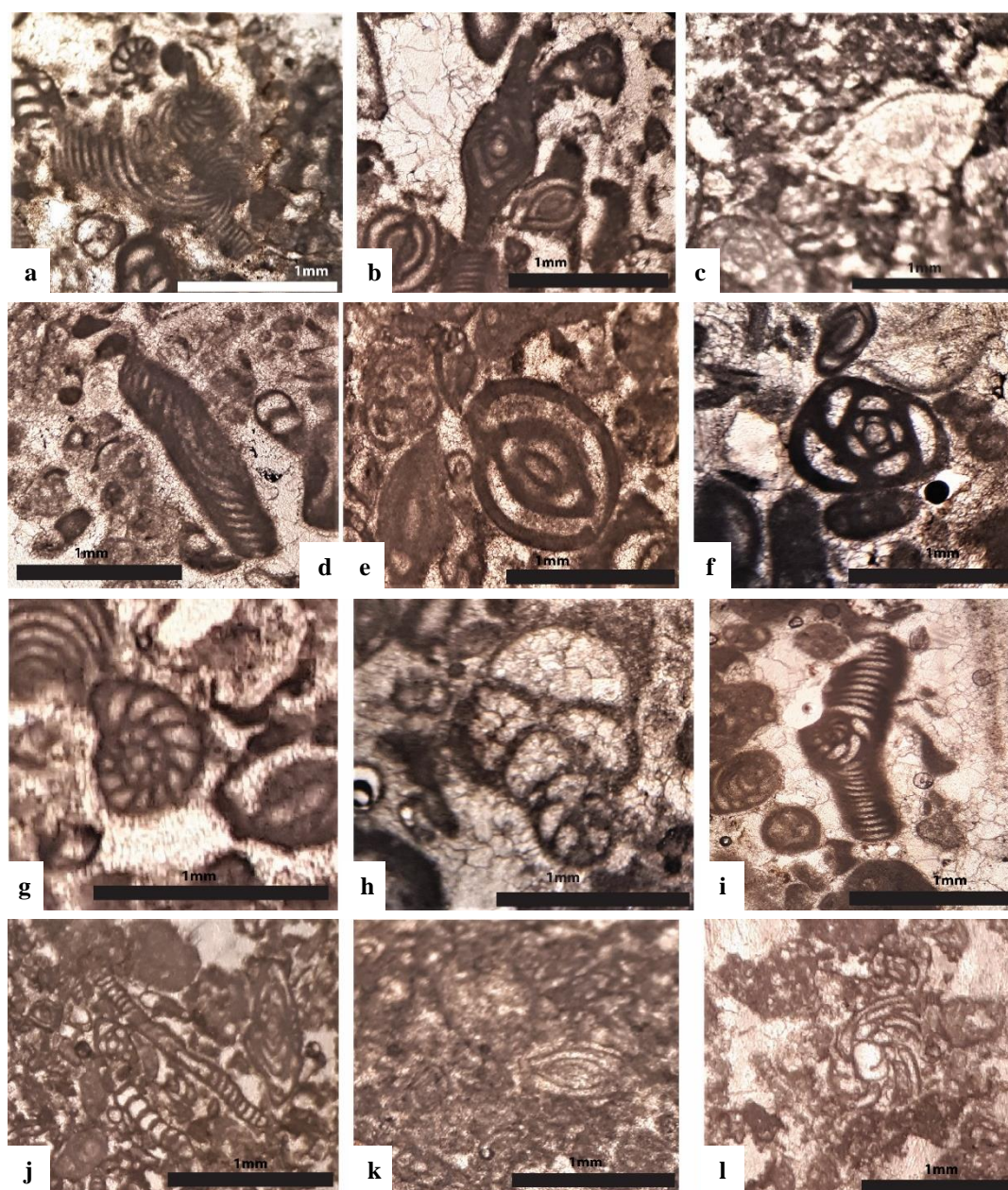


Fig.7: **a)** *Archaia hensoni* (Smout and Eames, 1958), oblique sections, sample no.1; **b)** *Archaia hensoni* (Smout and Eames, 1958), axial section, sample no. 1; **c)** *Neorotalia viennoti* (Greig, 1935), axial section, sample no.1; **d)** *Peneroplis evolutus* (Henson, 1950; in Hayward *et al.*, 2018), sample no.1; **e)** *Pyrgo* sp., sample no.1; **f)** *Quinqueloculin* sp., equatorial section, sample no.1; **g)** *Risananeiza pustulosa* (Boukhary and Abdelraouf, 2008) sample no. 1; **h)** *Textularia* sp. sample no.1; **i)** *Archaia hensoni* (Smout and Eames, 1958), transverse section, sample no.2; **j)** *Archias asmaricus* (Smout and Eames, 1958), sample no.2; **k)** *Nummulites vascus* (Joy and Leymerie, 1848 in Hayward *et al.*, 2018), axial section, sample no.2; **l)** *Heterostegina* sp., sample no.2

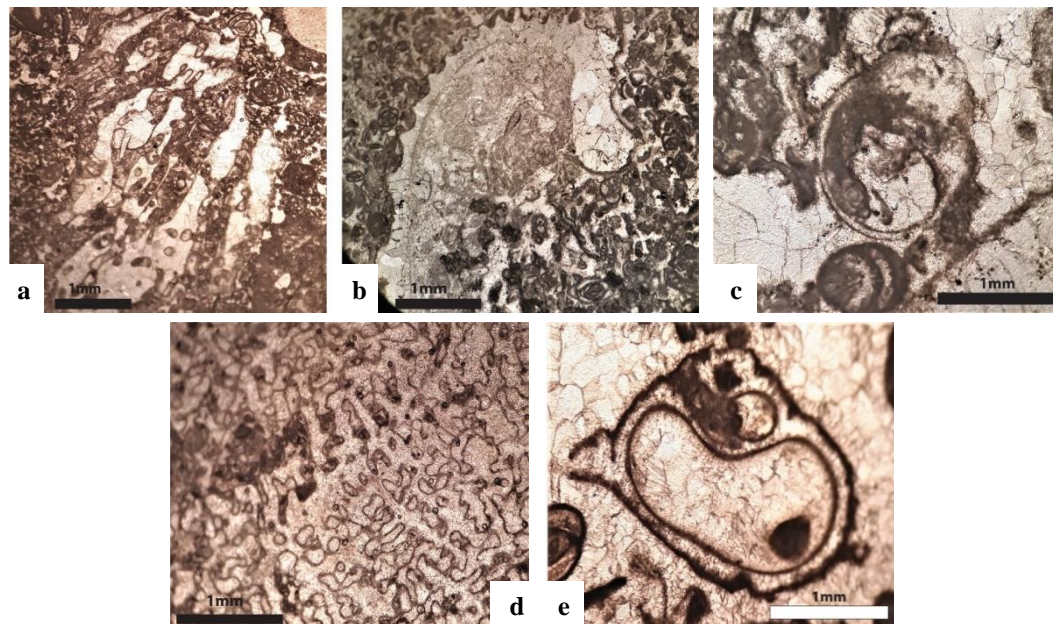


Fig.8: **a)** Bryozoan colonies, the zooecia wall structure has a fibrous microstructure, filled with sparry calcite cement, sample no.1; **b)** Brachiopod fragment has a fibrous crenulated shell. the fragments have been preserved, but the surrounding skeletal grains have been replaced with sparry calcite, sample no.1; **c)** Gastropoda shell, the gastropod has a geometric shape and the outer shell of the gastropod has been replaced with fine calcite cement, sample no.1; **d)** Coral, sample no.6; **e)** Pelecypoda (bivalves), the bivalves have micritized rims and the inner portion has been replaced with sparry calcite, sample no.3

▪ Paleoeological Interpretation

In this study the most significant paleoecological elements investigated are salinity, nutrients, and water depth. Paleoeological interpretation of the Oligocene carbonate platforms has been attempted by different authors (e.g., Hallock and Glenn, 1986; Pedley, 1998; Geel, 2000; Pomar, 2001; Romero *et al.*, 2002; Beavington-Penney and Racey, 2004; Pomar *et al.*, 2004; Pomar and Hallock, 2008; Brandano *et al.*, 2009, Van Buchem *et al.*, 2010; Flügel, 2010; Pomar *et al.*, 2012; and Taheri *et al.*, 2017}. In the Chattian succession of the studied section the most significant and dominant biogenic components are large, flat and perforate benthic foraminifera (*Operculina*, *Heterostegina*, *Spiroclypeus*, *Amphistegina*), coralline red algae and imperforate foraminifera (*Archaias*, *Peneroplis*, *Dendritina*). Other skeletal grains consist of, gastropod, bryozoan and bivalve fragments. Accordingly, the paleoecological conditions of the Anah Formation during the Chattian time could be interpreted as follows: Nummulitids and red algae thrived in oligophotic to mesophotic mid-ramp environments mostly at water depths that ranged from 40 to 80 m (Bosence, 1983; Hottinger, 1997; Beavington-Penney and Racey, 2004). Occurrence of large perforate benthic foraminifera (*Heterostegina*, *Neorotalia*, *Amphistegina*), coralline red algae, imperforate foraminifera (*Archaias*, *Peneroplis*, *Dendritina*, *Austrotrillina*), miliolids and molluscs (bivalves and gastropods) in the section indicate normal to higher sea water salinity in inner ramp setting (Flügel, 2010; Allahkarampour Dill *et al.*, 2012), (Taheri *et al.*, 2017), (Table 1).

Table 1: Vertical distribution of faunal composition, paleoecological elements, salinity, nutrient and depth estimation of the Anah Formation in the studied section

Age	Fossil Components	Seawater Salinity	Nutrients	Depth (m)
Chattian	Coralline red algae and large benthic foraminifera <i>Opeerculina</i> , <i>Heterostegina</i> , <i>Spiroclypeus</i> , <i>Amphistegina</i> , <i>Neorotalia</i> , <i>Archaias</i> , <i>Austrotrillina</i> , <i>Peneroplis</i> and <i>Dendritina</i> , with mollusca, coral, and miliolids,	Normal marine to slightly saline	Oligo-Mesotrophic	40 – 80

CONCLUSIONS

- The Anah Formation in the studied section at the Sangaw area is characterized by different microfossils of larger foraminifera, and skeletal fragments of coralline red algae and brachiopoda layer grain, coral, pelecypods-bivalves, gastropods and bryozoan shells.
- The Anah Formation is subdivided in this work into two assemblage biozones, that extend in age from Early to Late Chattian.
- The paleoecological study shows that the Anah Formation in the studied section was deposited under normal to higher seawater salinity, with water depth that ranged from 40 to 80 m, and under oligotrophic to mesotrophic nutrient conditions.

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